

Trajectoires planétaires et interplanétaires

Analyses orbitales autour de Mars et Vénus

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Ecole d'été du GRGS – Trajectoires planétaires et interplanétaires. September 5-9th 2016, Aussois, France.



- Planetary geodesy provides a valuable way to get constrain on the interior as well as on the thermosphere of planets.
- We shall present the basic concepts of space geodesy and detail application involving orbiting spacecraft :

Mars' gravity field and its seasonal variations Mars' deep interior and k_2 tidal Love number Mass of the Martian moons, Phobos & Deimos Venus' atmosphere density at high altitude







- size
- composition
- physical state
- temperature

- formation
- evolution
- dynamics
- life

Planetary atmospheres: study of 'school-case' for Earth's climate evolution



> Which processes drive atmospheric dynamics, climate, ... ?

Determination of present dynamics and past and future evolution of the atmosphere.

- To date, Geodesy (and astrometry as well) is the single way to constrain planetary interior (except for the Moon)
- Monitoring of the *gravity field,* rotation, tides, nutations.

 I will focus on the method, show the most recent results for Mars and Venus.



Aréophysique: Mars deep interior

Mars Global surveyor (2000)



Crustal magnetic anomalies on old terrain (4Gyr)



Mars had a Dipole Magnetic field 4 Gyr ago

What is the state, size, composition of Mars' core today?

Aréophysique: Mars' CO₂ seasonal cycle



> Up to 25% of total mass of the Martian atmosphere.

→ Seasonal gravity field variations

How to measure gravity field with a spacecraft?



Principle: Monitoring the free-fall motion of the spacecraft as for the falling-ball of an absolute gravimeter

Spherical and homogeneous planet: Keplerian trajectory



- The trajectories of planets around the Sun follow ellipses around the Sun (Kepler's laws).
- Theory of gravitation (Newton).
- The trajectory of an artificial satellite around a spherical and homogeneous planet follow Kepler's laws (first approximation)
- → The artificial satellite or orbiter trajectory follows an ellipse whose focus is at the center of mass of the planet. This elliptical trajectory remains in a plane fixed in space.

Positioning in space: 6 Orbital Elements (OE):



Ω and *i* allow positioning the orbital plane in space
 a and *e* provide the shape of the ellipse and *ω* allows positioning it in the orbital plane
 v or *M* (mean anomaly) allow positioning the orbiter on the ellipse at any given time *t*.

Spherical and homogeneous planet:



All Orbital Element do not change except the mean anomaly *M* → Kepler motion



Zonal terms : $J_I = -C_{I,0}$

All Orbital Elements vary → Perturbation of Keplerian motion

Representation of planet gravity field (1)

✓ Spherical harmonics expansion of gravity field at any point out of the planet

Central Potential *U*₀: (Point-mass representation) Keplerian motion

GM

With:

 (r, ϕ, λ) spherical coordinates of a point P outside the planet a_e equatorial radius of the planet

+GM

 $P_{lm}(sin(\phi))$ Legendre Polynomials

 $C_{\mbox{\scriptsize Im}}$ et $S_{\mbox{\scriptsize Im}}$ Stokes or harmonics coefficients

I and m degree and order of the spherical harmonics expansion Practically it is truncated at a maximal degree: *L.* $GM (=C_{00})$ Gravitational constant times Planet's mass. Perturbing potential U_p :

 $\sum_{n=l+1}^{\infty} \frac{a_e^{\prime}}{a_{e}^{l+1}} P_{lm}(\sin(\varphi))(C_{lm}\cos(m\lambda) + S_{lm}\sin(m\lambda))$

That potential is completely defined by the set of harmonics coefficients:

 C_{lm} , S_{lm} , as well as by a_e and GM.

> The precise reconstruction of the perturbed orbital motion of the spacecraft allows retrieving the harmonics coefficients: C_{lm} , S_{lm}

Representation of planet gravity field (2)

✓ W. Kaula has developed the perturbing gravitational potential U_p as a function of the Orbital Elements: *a*, *e*, *i*, ω , Ω *et M*.



$$S_{lmpq}(\omega, M, \Omega, \theta) = \begin{bmatrix} C_{lm} \\ -S_{lm} \end{bmatrix}_{l-m \ impair} \cos(\psi_{lmpq}) + \begin{bmatrix} S_{lm} \\ C_{lm} \end{bmatrix}_{l-m \ impair} \sin(\psi_{lmpq})$$

$$\psi_{lmpq} = (l - 2p + q)M + (l - 2p)\omega + m(\Omega - \theta)$$
 "gravity" argument of Kaula

 θ is the sideral angle of reference longitude at planetary surface (spin)

 This new expression of the gravitational potential allows deriving the variations of the Orbital Elements due to each harmonics coefficients C_{Im} et S_{Im}
 Lagrange's equations

Equation of motion: Lagrange's equations (central and perturbed gravitational potential)

$$\begin{aligned} \frac{da}{dt} &= \frac{2}{na} \frac{\partial U_P}{\partial M} \\ \frac{de}{dt} &= -\frac{\sqrt{1-e^2}}{na^2 e} \frac{\partial U_P}{\partial \omega} + \frac{1-e^2}{na^2 e} \frac{\partial U_P}{\partial M} \\ \frac{di}{dt} &= \frac{-1}{na^2 \sqrt{1-e^2} \sin i} \frac{\partial U_P}{\partial \Omega} + \frac{\cos i}{na^2 \sqrt{1-e^2} \sin i} \frac{\partial U_P}{\partial \omega} \\ \frac{d\Omega}{dt} &= \frac{1}{na^2 \sqrt{1-e^2} \sin i} \frac{\partial U_P}{\partial i} \\ \frac{d\omega}{dt} &= \frac{\sqrt{1-e^2}}{na^2 e} \frac{\partial U_P}{\partial e} - \frac{\cos i}{na^2 \sqrt{1-e^2} \sin i} \frac{\partial U_P}{\partial i} \\ \frac{dM}{dt} &= n + \frac{2}{na} \frac{\partial U_P}{\partial a} - \frac{1-e^2}{na^2 e} \frac{\partial U_P}{\partial e} \end{aligned}$$
Unperturbed Keplerian motion: $Up=0$ therefore: $M=n(t-t_p)$

- Lagrange's equations for a motion driven by forces derived from a potential.
- \succ '*a*, *e*,*i*,Ω,ω, et *M*' are the Orbital Elements (OE)
- 'n' is the mean motion of the unperturbed Keplerian motion.
- 'Up' is the perturbing potential accounting for the non-spherical terms of the gravititaional potential of the planet.

- > Perturbed Keplerian motion: $Up \neq 0$
- One cannot obtain accurate solution of Lagrange's equations using the perturbation method. Hereafter, is presented the Kaula's method (1966).

Kaula's solution (1966)



Integration of Lagrange's equation into two steps:
 1) Secular drifts then 2) Périodic terms

First order perturbation: Precessing Orbit



 Secular drift of the ascending node around the rotation axis of the planet:

$$\frac{d\Omega}{dt} = \frac{3n\cos(i)}{2(1-e^2)^2} \left(\frac{a_e}{a}\right)^2 C_{20}$$

Proportional to C₂₀ and cos(*i*)

First order perturbation: Precessing Orbit Drift of percienter argument: ω



• Secular drift of the pericenter argument (rotation of the ellipse in the orbital plane:

$$\frac{d\omega}{dt} = -\frac{3n}{(1-e^2)^2} \left(\frac{a_e}{a}\right)^2 \left(1-\frac{5}{4}\sin^2(i)\right) C_{20}$$

•Proportional to C₂₀

Longest period perturbations: Oscillation of pericenter position and excentricity

Again from Kaula's linear theory:

$$\frac{dp}{dt} + j\dot{\omega}_{s}p = -\frac{3n}{2(1-e^{2})^{3}}\sin(i)\left(\frac{a_{e}}{a}\right)^{3} \left[\left(\frac{5}{4}\sin^{2}(i)-1\right) + \left(\frac{1}{2\sin^{2}(i)}-\frac{31}{8}+\frac{15}{4}\sin^{2}(i)\right)e^{2}\right]C_{30}$$

Important remarks:

(1) " $p=p_0 \exp(-j\omega)$ " is the excentricity vector. It follows periodic variations: $T_s = \frac{2\pi}{\dot{\omega}_s}$ The period depends on the secular drift of pericenter position (i.e. a, e, i, and C₂₀)

(2) The amplitude of this periodic variation varie as $(1/a)^3$ so less than as $(1/a)^2$

(3) This amplitude is proportional to sin(i). So it is zero for equatorial orbit, $i=0^{\circ}$ and maximal for polar orbits, $i=90^{\circ}$

Others periodic perturbations

$$\psi_{lmpq} = (l - 2p + q)M_s + (l - 2p)\omega_s + m(\Omega_s - \theta) \neq 0$$

> The gravitational potential depends on all Orbitla Elements (OE) and is expressed through the summation over all the indexes *I*,*m*,*p*,*q* of Ψ_{Impg}



 > Important result: The amplitude of periodic variations is proportional to 1/ψ_{lmpq} Short periodic perturbations are lower than long period perturbations.
 > Resoances: ψ_{lmpq} = 0. Kaula's approach can only allow detecting those resonances for any orbit (*a*,*e*,*i*) for each associated harmonics coefficients c_{lm} et s_{lm}.

Orbit filters gravity field through resonances



Kaula's solution
 → Orbital velocity perturbation on radial direction due to each degree & order: C_{Im} et S_{Im}

- Comparison with error on velocity measurements --> Resolution of gravity field (SNR=1)
- Example: MGS/ODY \rightarrow Maximal degree Lmax: 60.

First idea of gravity field resolution from a given orbit.

Mean OE: a=400 km, e=0.01, i=92.9°

Marty et al., 2009

From theory ... to practice

- Kaula's approach allows studying the sensitivity of a given orbit to harmonics of gravity field C_{lm} et S_{lm} (i.e. given *a*,*e*,*i*).
- Nevertheless, it cannot allow (as for any other analytical approach) to obtain an accurate orbit reconstruction *(perturbation method, non-conservative forces, ...).*
- *In practice*, one numerically solve for the fundamental equation of dynamics.
- The gravity field harmonics are then estimated through the fit of an initial gravity field to **tracking data** of the orbiting spacecraft.
- The non-conservative forces must also be estimated, because they are not well known. This is the major limitation of the method.

Precise determination of the spacecraft orbital perturbations



Tracking data: Radio-link between the Earth and spacecraft (generally 2-way X band)

Precise Orbit Determination (POD) GINS⁽¹⁾: orbitography software:

- Calculation of orbit from least squares fitting of a dynamical model of motion to tracking data
- Determination of the planetary gravity field (including its time variations), k₂ tidal Love number, *GM* of the moons, ...

(1) Géodésie par Intégration Numériques simultanées, developed by the French space agency (CNES) and further adapted to planetary geodesy applications by the Royal Observatory of Belgium (ROB).

NASA's DSN and ESA's ESTRACK networks of tracking stations

> To track any spacecraft or Lander in the solar system



Diameter of 25 to 70 meters.

Precision: 0.02-0.05 mm/s (1-3 mHz) < ~1 m. But bias of ~3-4 m. Dating: Maser clocks

- Radio-link (transmitting & receiving) for data & telemetry
- Doppler & range radio-tracking mainly at frequencies of 8.4 Ghz (X band)

Application of Doppler effect to spacecraft velocity determination

 Doppler effect: range-rate variations relative velocity: Earth-spacecraft



Important remark: Doppler effect = 0 when the Line-Of-Sight direction is perpendicular to the spacecraft velocity direction

Doppler signal variations Earth-Mars radio-link



Precise orbit determination: in some words

- To model all the forces acting on the spacecraft and to numerically integrate its motion (*not only gravity*)
- To calculate the associated predicted tracking data
- To perform an iterative least squares process on the radio-science tracking data (Doppler and ranging) in order to adjust parameters related to the data and to *dynamical model of spacecraft motion*
- We use a dedicated software *GINS*¹. Two others software are able to do same calculations *DPODP* at JPL (USA) and *GEODYN* at GSFC (USA). Others operational or in development in Europe.

¹ Géodésie par Intégrations Numériques Simultanées, developped by CNES and applied by ROB for planetary geodesy applications

Martian spacecraft tracking data



Spacecraft dynamical model

> Model of all the forces acting on the spacecraft

- Gravitational forces:
 - Mars' static gravity field (JGM95J from JPL)
 - Point mass representation of other solar system bodies using JPL planetary and Martian moon ephemerides.
- Non-gravitational forces :
 - Atmospheric drag (atm. Density model)
 - solar pressure radiation, albedo & IR radiations.
 - Residual accelerations induced by each unbalanced wheel off-loading (WoL) or angular moment desaturation event.

Most unknown forces to be adjusted:

Drag (specially MEX pericenter altitude at 250 km) and WoL accelerations. Additional difficulty for MEX: *Non-continuous tracking*.

Non-gravitational accelerations







Spacecraft as flat plates (6 for the bus & 4 for solar panels) with known area and optical properties (reflectivity and absorptivity).

The orientation of this model in space is given by quaternions (bus, ...).

Shift between center of mass and the HGA center of phase. Mass history.

Example of quaternions of the bus of MEX



Attitude mode : Earth or Mars pointing or fixed inertial modes.

Precise Orbit Determination: POD



The quality of the estimated geophysical parameter strongly rely on the precision of the reconstructed orbit of the spacecraft.

Non-continuous tracking of MEX. A problem !



- Depending on the data-arc, successive MEX revolutions around Mars are not tracked at all.
- Difficulty to correctly determine the effect of the desaturation (WoL) and of the drag at each pericenter pass (altitude of 250-700 km).

Post-fit residuals of tracking data: MEX



- Good residuals in Doppler and range. Effect of interplanetary plasma on signal at solar conjunction (September 2004 and October 2006).
- > Test of orbit quality: Orbit overlap (test of internal coherency).

Precision on MEX & Odyssey reconstructed orbit



- MEX orbit precision is about ten times worse than Odyssey (MGS) orbit precision.
- This is mainly due to the lack of tracking data at MEX pericenter which does not permit to precisely estimate the atmospheric drag orbital perturbations
Estimation of the atmospheric drag effect: Drag scale factor $F_D(1)$



Very poor tracking coverage at MEX pericenter passes occurs during repeated period of several weeks.

>This limits the orbit accuracy in the along-track direction.

Estimation of the atmospheric drag effect: Drag scale factor $F_D(2)$



> Solar UV flux minimum decreases the density of Mars high atmosphere.

- Thus, this decreases the drag force onto MEX, and makes more difficult to estimate the drag scale factor for this period.
- Similar behavior for MGS and ODY spacecraft

Estimation of the solar pressure scale factor: F_s



Solar pressure factor estimate for MGS and ODY

 F_r : Reconstructed force F_m : Modeled force

$$\vec{F}_r = F_S \vec{F}_m$$

- Increasing MEX solar pressure scale factor with time.
- Similar behavior for MGS between 1999-2002.
- Darkening of the faces of the spacecraft with time?

Corrections to the WoL residual accelerations: MEX



Accelerations associated with WoL events not well resolved along the directions of the X and Y axes of the spacecraft frame.

This degrades the determination of MEX orbit in directions normal to the along track direction.

Effect of attitude maneuvers on the orbit



• A few meters at each attitude maneuver event (1 mm/s of Delta_Velocity at eahc maneuver in 3 directions

Spacecraft range tracking and planetary ephemerides



The range tracking data allows planets ephemerides to be updated. Especially for Mars.

Current Mars' static gravity field solutions

GRGS/ROB solution, Marty et al., 2009

Spatial resolution about 300 km







Limitation due to the orbit sensitivity to Doppler tracking precision and to non-gravitational forces: **Degree strength around 70.**

Challenge: To detect fine temporal variations of the first zonal harmonics and the k_2 tidal Love number

Degree strength of gravity field solution

$$U = \frac{GM}{r} + GM \sum_{l=1}^{l=\infty} \sum_{m=0}^{m=l} \frac{a_{e}^{l}}{r^{l+1}} P_{lm}(\sin(\varphi))(C_{lm}\cos(m\lambda) + S_{lm}\sin(m\lambda))$$

Intensité du potentiel en 1/r^{l+1}

→ Sensibilité à un degré maximal *L* donné par l'altitude au périgée h_p :



- ✓ Kaula's rule of thumb provides a good approximation of the degree strength "L" of the gravity field solution.
- ✓ Elliptical orbit offers lower altitude than circular orbits but only at pericenter, thus generally doesn't cover the entire planetary surface (like for MEX, Magellan).

Degree strength variations: Venus gravity field solution



Anderson & Smrekar, 2006

Gravity field variations and Mars' CO₂ seasonal mass budget



Seasonal gravity field to determine mass transfer budget, but insufficient precision to constrain the models of CO_2 seasonal deposits (*Karatekin et al., JGR, 2006*).



GINS simulations (Rosenblatt et al., AGU fall meeting 2005)



Mars' atmosphere CO₂ seasonal cycle

MEX + MGS can improve the solution of first zonal harmonics variations, thus the seasonal mass budget, given the orbits are well resolved.

We need to perform an accurate spacecraft orbit. Yes, but how much accurate?

Current solutions of Mars' C₃₀ seasonal variations from MGS/ODY tracking data

GRGS/ROB solution (Marty et al. 2009)



JPL solution (Konopliv et al., 2006)



- C₃₀ seasonal variations determined from MGS and ODY tracking data.
- Slight differences in amplitude between JPL (Konopliv), GSFC (Lemoine) and GRGS/ROB solutions.

Current solutions of Mars' C₂₀ seasonal variations from MGS/ODY tracking data





- GINS procedure: Fit of spacecraft simulated positions computed with time variable gravity on spacecraft simulated positions computed without time variable gravity on successive 4-days over 1 Martian year.
- MGS orbit (actual orbit accuracy is 1-2 meters on average): Signature of odd zonal harmonics < 1 meter & even zonal harmonics < 10 cm</p>
- MEX orbit (actual orbit accuracy is 20 meters on average): Signature of time variable zonal harmonics ~10 cm !

Fit of time variable C_{20} & C_{30} from MGS/ODY and from MEX



Effect of poor MEX tracking coverage on time variable C_{20} & C_{30} solutions



- The formal uncertainty on C₂₀ & C₃₀ solutions from MEX are systematically higher when tracking data are not available at pericenter passes, i.e. when the drag cannot be well resolved !
- This limits our ability to retrieve the information about the time variable gravity from MEX orbits

How to improve Mars seasonal gravity solutions? (1)

Odd zonal harmonics variations detection favored by polar orbit:

$$\frac{dp}{dt} + j\dot{\omega}_{s}p = -\frac{3n}{2(1-e^{2})^{3}}\sin(i)\left(\frac{a_{e}}{a}\right)^{3} \left[\left(\frac{5}{4}\sin^{2}(i)-1\right) + \left(\frac{1}{2\sin^{2}(i)}-\frac{31}{8}+\frac{15}{4}\sin^{2}(i)\right)e^{2}\right]C_{30}$$

Even zonal harmonics variations detection favored by non- polar orbit:

$$\frac{d\Omega}{dt} = \frac{3n\cos(i)}{2(1-e^2)^2} \left(\frac{a_e}{a}\right)^2 C_{20}$$

New tracking data of orbiter with non-polar orbit are welcome Opportunity: ESA's Exomars spacecraft (**Trace Gas Orbiter**) TGO Non-polar (**74**° **inclination**) near-circular orbit (400x400 km).

How to improve Mars seasonal gravity solutions? (2)



Challenging !

Tidal potential (*k* Love numbers) and the state and size of Mars' core



Current knowledge of the Martian core from geodetic data (e.g. Yoder et al., 2003; Konopliv et al., 2006; Marty et al., 2009)





Core radius estimates given possible mantle temperature end-members, mantle rheology, and crust density and thickness range (from A. Rivoldini, ROB).

All solutions of k_2 indicate a liquid core inside Mars ($k_2 > 0.08$), but discrepancies still remain, thus implying a too large uncertainty for core radius estimate (+/- 250 km) to better constrain Mars' deep interior structure.

This is due to the difficulty to extract the weak signal of the k_2 from the current reconstructed orbits of the Martian spacecraft.

Secular effect of the k₂ Love number on orbital perturbations

$$U_{n,m,p,p',q,q'} = \frac{GM_P}{a_P} \left(\frac{R_0}{a_P}\right)^n \left(\frac{R_0}{a}\right)^{n+1} k_n \delta_m \frac{(n-m)!}{(n+m)!} F_{n,m,p'}(i) F_{n,m,p'}(i_P)$$

$$G_{n,p,q}(e) G_{n,p',q'}(e_P) \cos \psi_{n,m,p,p',q,q'}$$

 $\psi_{n,m,p,p',q,q'} = (n-2p+q)M - (n-2p'+q')M_P + (n-2p)\omega - (n-2p)\omega_P + m(\Omega - \Omega_P)$



Analytical developments, following Kaula's approach: near-polar orbits are not the most suitable orbits for k₂ determination

Improved determination of the mass of the Martian moons with MEX.



• A small secular drift of spacecraft orbits is caused by the Martian moon masses

•The higher elliptical orbit of MEX makes it more sensitive to the gravitational attraction of the Martian moons.

• Secular solution of the mass of the moons from stacking together all MEX data-arcs.

• Flyby solution with MEX too.

 Taking advantage of the higher elliptical orbit of MEX to improve the mass of the Martian moons: secular & flyby (Phobos) solutions.

MEX orbital motion perturbations at close flyby to Phobos



Much better determination of Phobos mass



Determination of the gravity field of Phobos



MEX/HRSC image

Results using flyby with Mars Express

Flyby distance and Doppler geometry:



 Closest approach ever: 58.71 km (from mar097 ephemeris) 58.61 km (from New IMCCE ephemeris)
 58.77 km (from mar085 ephemeris) 59.21 km (from Lainey et al. (2007) epehemris)



Quasi-continuous tracking up to 2 orbital revolutions before and after the flyby.
 POD fit using all this tracking data: long data-arc of about 1.2 days



✓ But irrealistic scale factor (drag, solar pressure) Not shown here



 \checkmark Quasi-continuous tracking up to 2 orbital revolutions before and after the flyby.



- ✓ Drag scale factor: 0.77 +/- 0.14 & Solar pressure scale factor: 1.1743 /- 0.0016
- ✓ Independent of Phobos' ephemeris.
- ✓ Taken as a priori value and constraints for GM/C20 fit with data-arcs including flyby



✓ Data-arc length: 3, 2 and 1 orbital revolutions centered on flyby

Estimating dynamical Phobos' gravity field from flyby



- ✓ GM estimated close to initial value of 0.711 E+06 m3/s2 with formal error of about 0.02%.
- ✓ C₂₀ estimated close to about -0.32 with formal error of about 0.002 (0.6%).
- Precise solution of GM & C₂₀ but large biais to C₂₀ solution (physically unplausible)
 'Slight' biais between solutions using mar097 vs ESPaCE-IMCCE ephemeris

Possible explanation \rightarrow biased solutions due to error on ephemeris?

Simulation of Doppler tracking data:

Simulation process

Data-arc duration : 1 revo., 2 revo., 3 revo., around flyby Ground station as for MEX tracking of December 29th 2013 Initial state vector (position/velocity) from FD MEX orbit *Initial GM=0.711 10⁶ m³/s² and C*₂₀=-0.1

Simulated Doppler data (60sec sampling time) with white noise at 0.02 mm/s.

Modified parameter value as *a priori* value: $GM = 0.709 \ 10^6 \ m^3/s^2$ and C20 = 0.0 and perturbed ephemeris at 1000 meters, 100 meters and 10 meters level.

Fitted parameter: Initial state vector (position/velocity) of MEX at the beginning of the data-arc and GMand C_{20} .

Results: Solution of *GM* and C20 (adjusted value and formal error). Impact of ephemeris error

Simulation of Phobos' ephemeris error



 Simulation of Phobos' ephemeris error: Shifting Phobos position with a constant biais of 1km.



Result of simulations: effect of Phobos' ephemeris error

✓ Slight bias on GM retrieval: 0.1%

✓ Large bias on C_{20} retrieval

✓ Phobos ' ephemeris error of 1 km mimics the bias observed on true data for the C₂₀ solution (*better simulation adding the C*₂₂ to be done)

✓ 100 meters on ephemeris bias → about 10% of biais on C_{20} 10 meters on ephemeris bias → about 1% of bias on C_{20}

Summary for Mars

- ✓ Reconstruction of spacecraft orbiting motion is currently the single way to determine the gravity field of planets (solution as, C_{lm} , S_{lm})
- ✓ Each orbit provides its own sensitivity to each C_{lm} , S_{lm} (spatial resolution, "degree strength").
 - → Combination of spacecraft with different orbits to improve gravity field solution.
- ✓ The precision on gravity field solution directly depends on precision on orbit reconstruction (→ Mars seasonal gravity variations, k_2 and Phobos gravity field)
- ✓ Single straightforward observation of liquid core inside Mars & Venus → deep interior structure and evolution.

The state-of-the-art: The GRAIL mission







NASA's mission GRAIL: Two spacecraft orbiting the Moon. Radio-link between both.

Best gravity field of the Moon ever obtained And the best gravity field of the Solar system to date: Spatial resolution ~ 25 km *Four times better than GOCE Earth's solution*

... for Mars ? GRAIL-like, GOCE-like experiment?
The state-of-the-art: The GOCE mission







ESA's mission GOCE: Best gravity field of the Earth ever obtained Spatial resolution 100x100 km Precision on the Geoid: 1 cm GINS used to produce this new solution



Density measurements of upper atmosphere of planets using Doppler tracking data of orbiting spacecraft:

Recent results with Venus Express



Some basic physics: 'The drag paradox'

Drag force is opposite to velocity

> We may guess a deceleration of the spacecraft

 $\delta W = \vec{F}_D \cdot \vec{V} dt = -F_D V dt$

 \succ Actually it corresponds to an acceleration !

➢ Drag force is a dissipative force that decreases the mechanical energy → <u>dE < 0</u>:
➢ How does it affect a circular orbit?

(1) Decrease of semi-major axis "a": $\delta W = dE = d\left(-\frac{GM}{2a}\right) = \frac{GM}{a^2} da$ (2) Acceleration of orbital speed "V": $dV = d\left(\sqrt{\frac{GM}{a}}\right) = -\frac{1}{2}\sqrt{\frac{GM}{a^3}} da$ (3) Shortening of orbital period "T":

$$dT = d\left(\frac{2\pi}{\sqrt{GM}}a^{3/2}\right) = 3\sqrt{\frac{a}{GM}}da$$

$$\blacktriangleright dE < 0$$
, hence $da < 0$

$$\blacktriangleright da < 0$$
, hence $dV > 0$

$$\succ$$
 da < 0 hence dT < 0

Motivation for VExADE

 Existing density models built with orbital period approach and also mass spectrometers, probes, ground-based observations ...):

VTS3 (Hedin et al., 1983) VIRA (Keating et al., 1985) Venus Gram (recently released)

fortran-routine look-up table fortran-routines

• These models are constrained for equatorial latitudes

 \rightarrow Needs of new data for higher latitudes

• Venus Express Atmospheric Drag Experiment: VExADE

To probe *in situ* the polar thermosphere at low solar activity

The VExADE campaigns



- 10 campaigns have been performed, each over a few successive days from 2008 to 2012.
 Back to nominal orbit (peri at 250 km) between each campaign.
- Derivation of Venus upper atmosphere neutral density using Doppler tracking data \rightarrow Precise Orbit Determination (POD).





GINS: Géodésie par Intégrations Numériques Simultanées, developed by CNES and applied by ROB for planetary geodesy applications

Venus Express dynamical model

- > Model of all the forces acting on the spacecraft
- Gravitational forces:
 - Venus' static gravity field (JPL 180x180)
 - Point mass representation of other solar system bodies using updated JPL planetary ephemeris.
- Non-gravitational forces :
 - Atmospheric drag (atm. Density model): scale factor FD
 - solar pressure radiation.
 - albedo & IR radiations (mean value).
 - Residual accelerations induced by each unbalanced inertial Wheel-Off-Loading (WoL)

Venus Express "Box-and-Wings" model



✓ "Box-and-Wings" model or "Macro-Model" (from navigation team):

1) Shape of spacecraft:

flat plates (6 for the bus & 4 for solar panels) with known area

2) Optical properties:

known reflectivity and absorptivity coefficients

3) Orientation of this Box-and-Wings model in space (quaternions)

The shift between the s/c center of mass and the HGA1 or HGA2 center of phase

Precise modeling of the drag acceleration

$$A_{D} = -\frac{1}{2} \rho F_{D} \sum_{i=1}^{i=N} C_{Di} \frac{S_{i}}{m} (V_{r}.n_{i}) V_{r}$$

Where:

 $\begin{aligned} A_d &= Drag \ acceleration \\ \pmb{\rho} &= \pmb{Predicted atmospheric density} \ (\textit{Hedin-VTS3 model or Venus-Gram}) \\ S_i &= surface \ area \ of \ the \ plate \ \ll \ i \ \gg C_{Di} = Drag \ coefficient \ of \ the \ plate \ \ll \ i \ \gg V_r = velocity \ vector \ of \ spacecraft \ w.r.t. \ atmosphere \\ n_i &= unit \ vector \ normal \ to \ the \ face \ \ll \ i \ \gg (orientation \ of \ the \ plate) \\ m &= spacecraft \ mass \\ \pmb{F_D} &= estimated \ drag \ scale \ factor \ (a \ priori = 1, \ no \ a \ priori \ constraint) \\ N &= number \ of \ flat \ plates \ of \ the \ macro-model \end{aligned}$

$$\rho_{measured} = F_{D}.\rho_{predicted}$$

✓ The fit of the force model to the tracking data provides a drag scale factor of the predicted (modeled) atmospheric density model.

Ground segment: Tracking stations

>To track any spacecraft or Lander in the solar system



Diameter of 25 to 70 meters. Precision: 0.02-0.05 mm/s (1-3 mHz) < ~1 m. But bias of ~3-4 m. Dating: Maser clocks Radio-link for data & telemetry

 Doppler & range radio-tracking mainly at frequencies of 8.4 Ghz (X band) (sometimes dual-frequency X/S bands to correct ionospheric and interplanetary plasma perturbations)

Example of VExADE good tracking data coverage



- 4 hours of tracking performed by ESTRACK and DSN around each pericenter pass and 8 hours at high altitude performed by ESTRACK.
- Data-arcs for POD are about 1 day of duration: From one near-apocenter to the next to avoid perturbing WoL maneuvers.

Example of Doppler post-fit residuals (VExADE campaign #2 – Pericenter altitude 177 km)

VEX ADE-2: 2009/10/17

VEX ADE-2: 2009/10/17



Good post-fit residuals: rms value of ~ 8.3 mHz (or 0.3 mm/s).

It does not necessarily mean good estimates of drag sale factor.

Example of POD results: VExADE campaign#2



Other parameters than drag scale factor must have correct values
Significant variations drag scale factor can occur (true signal?)

Example of Doppler post-fit residuals (VExADE campaign #2 – Pericenter altitude 177 km)

VEX ADE-2: 2009/10/17



 Doppler signatures at pericenter pass due to error on other forces: The gravity field ?.

Realistic error on density estimate from POD



- Observed Doppler signature at pericenter can be reproduced by simulating the effect of long wavelength gravity field error.
- The drag scale factor estimate has a bias at around 6 times its formal error.
- Density estimate error had to be scaled by a factor of 6 (5% \rightarrow 30%).

Compiling VExADE

campaign#1 to #10

VExADE drag campaign#1 to #10: Dataset



80° -90°, local time: 1h, 7h, 18h Solar activity: F10.7 ~66-145.

VExADE campaigns#1 to #10: Results



- Venus' upper atmosphere density at polar areas is about half the predicted density by model (Hedin-VTS3).
- Checked by independent method using the inertial wheels of the spacecraft.

Day-to-day variability of observed densities



- Day-to-day variability in 1 km altitude range (at low altitude) not accounted for by the VTS3 model.
- Thermospheric temperature variations? Dynamics effects? Others?
- Torque measurements may help to decipher that observation.



Torque measurements

 $T = \frac{1}{2} C_D \rho A_{eff} V^2 (r_{SA-cop} - r_{SC-com})$



- ✓ Special attitude mode of spacecraft at VExADE pericenter pass.
- Torque induced by air pressure measured by rotation rate variati of inertial wheels.
- ✓ Independent measurement of density.

Profile of density along the pericenter pass



Figure 3: Torque around SC x-axis plot versus altitude. The pre-pericentre and postpericentre is shown in black and red respectively. On 2010-02-25 the SC was flying from night to day, on 2010-04-14 the SC was flying from day-to-night. For the 2010-10-18 and 2010-10-20 plots the SC is flying almost along the terminator.

A new model of Venus' polar upper atmosphere density.



- The VExADE estimated densities have contributed for altitudes above 165 km
- Additional remote sensing data from VEX SPICAV-SOIR instrument and from Earth's based JCMT sub-millimeter telescope have contributed at lower altitude range of 100-130 km.

- This new model has been used by the ESOC/ESA teams for the preparation of the July 2014 aero-braking phase of the VEX mission.
- The spacecraft system accelerometer has been used to probe density, still at polar latitudes but at altitudes as low as 140 km (*Nature Physics, 2016*)



From NASA

Summary for drag

- ✓ Precise orbit reconstruction allows probing *in-situ* the density of the planetary atmosphere at altitude of pericenter.
- ✓ Accuracy limited by accuracy on gravity field. Limited sampling (VEX)
- ✓ Application to Mars: Circular orbits (MGS, ODY, update Bruinsma et al., 2014, MRO) and elliptical orbits (MEX, MAVEN)

